

# A Spatial Bayesian Hierarchical Model for a Precipitation Return Levels Map

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**NCAR**



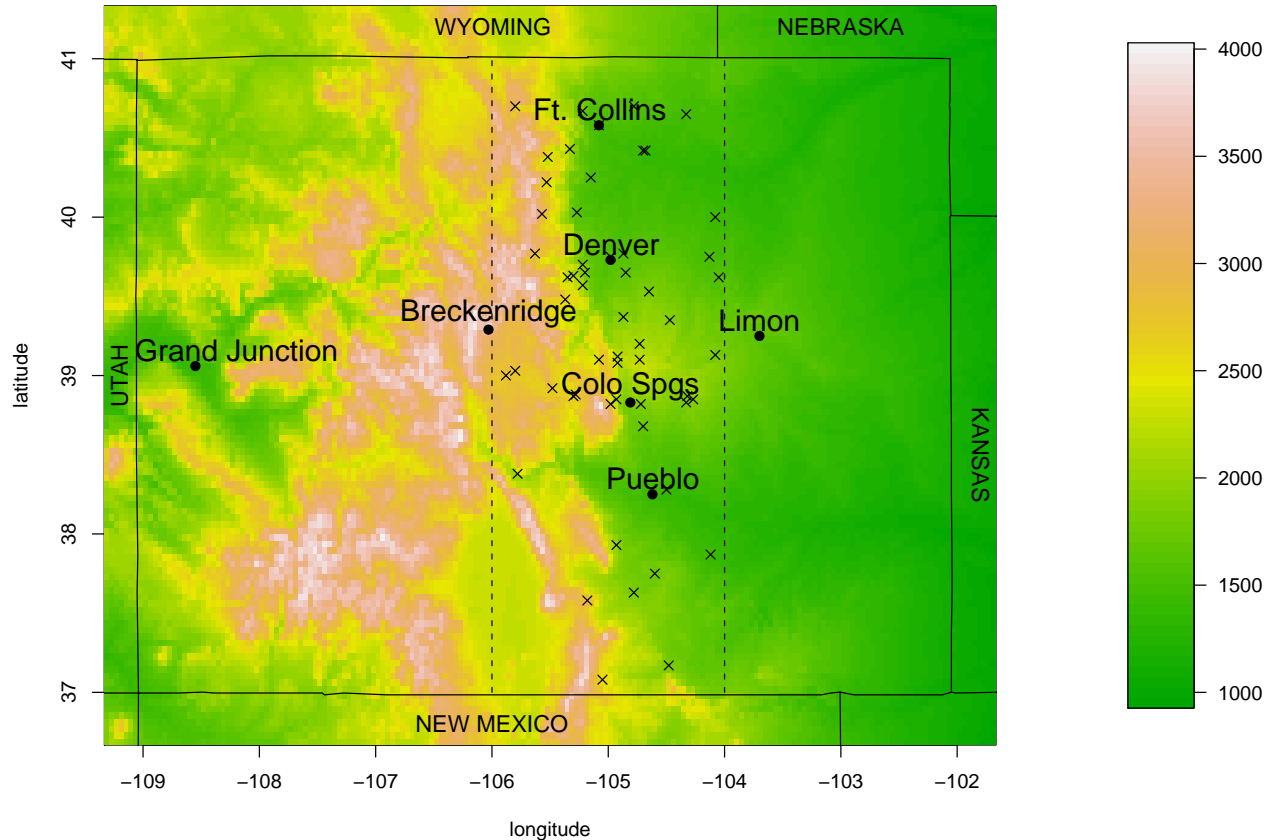
# Project Background

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*Goal:* To produce a map which describes potential extreme precipitation for Colorado's Front Range.

- Part of a larger NCAR project on flooding
- 1973 NOAA/NWS Precipitation Atlas is currently used
  - no uncertainty estimates
  - outdated extremes techniques
  - 30 more years of data
- Current NWS effort to produce updated maps
  - maps produced for two US regions (not Colorado)
  - using RFA methodology of Hoskings and Wallis
- Precipitation atlases provide **return levels** measures

# Study Region



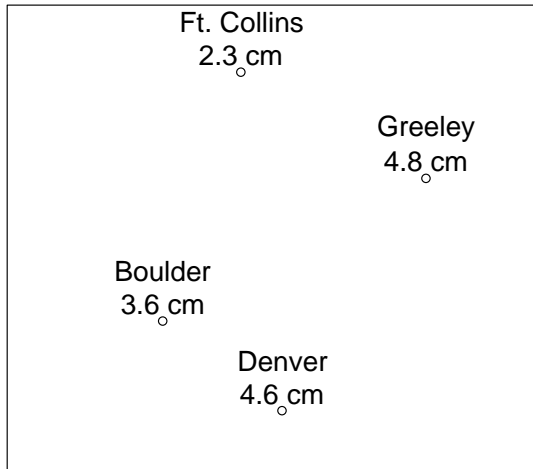
Data: 56 weather stations, 12-53 years of data/station, Apr 1 - Oct 31, 24 hour precipitation measurements

# Weather, Climate and Spatial Extremes

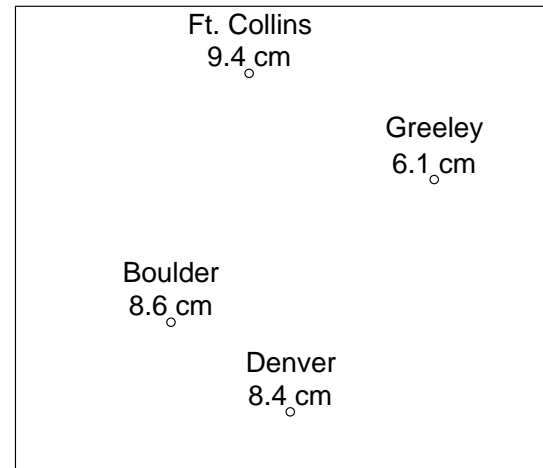
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- Modeling observations.
- Characterizing spatial dependence in the **data**.
- Short-range dependence.
- Modeling return levels.
- Characterizing dependence in the **distributions**.
- Longer-range dependence.

Max Daily Prcp 2000



25 Year Return Level



# Modeling Climatological Extremes

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- We use a POT approach, and assume exceedances over a threshold  $u$  are described by a  $GPD(\sigma_u, \xi)$ .
- We assume the climatological extreme precipitation is characterized by a latent process –  $\sigma_u$  and  $\xi$  are functions of location
- Return levels:

$$z_r(\mathbf{x}) = u(\mathbf{x}) + \frac{\sigma_u(\mathbf{x})}{\xi(\mathbf{x})} \left[ (r n_y \zeta_u(\mathbf{x}))^{\xi(\mathbf{x})} - 1 \right].$$

- $n_y$  is the number of observations in a year.
- $\zeta_u(\mathbf{x})$  is the probability an observation exceeds  $u$ .

# Model Goals

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- Utilize extreme value theory (GPD)
- Pool the data from the stations into one model – different from RFA
- Model should have spatial coherence
- Should utilize available covariates – elevation and mean Apr-Oct precipitation
- Should be flexible enough to be able to compare models
- Produce measures of uncertainty

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## *Spatial Bayesian Hierarchical Models:*

Independent 3-layer (data, process, prior) models for threshold exceedances and exceedance rates.

# Exceedances Model - Data Level

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Let  $Z_j(x_i)$  be the precipitation amount recorded at the station located at  $x_i$  on day  $j$ . We assume that precipitation events  $Z_j(x_i)$  which exceed a threshold  $u = .45$  inches are GPD, whose parameters depend on the station's location.

$$P\{Z_j(x_i) - u > z | Z_j(x_i) > u\} = \left(1 + \frac{\xi(x_i)z}{\exp \phi(x_i)}\right)^{-1/\xi(x_i)}$$

# Exceedances Model - Process Level

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$\phi(\mathbf{x})$ : Modeled with standard geophysical methods  $\rightarrow$  Gaussian process

$$\begin{aligned}\mu_\phi(\mathbf{x}) &= f(\boldsymbol{\alpha}_\phi, \text{covariates}(\mathbf{x})) \\ &= \alpha_{\phi,0} + \alpha_{\phi,1}(\text{elevation}) \text{ (for example)}\end{aligned}$$

$$k_\phi(\mathbf{x}, \mathbf{x}') = \beta_{\phi,0} * \exp(-\beta_{\phi,1} * \|\mathbf{x} - \mathbf{x}'\|_2)$$

# Exceedances Model - Process Level

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$\xi(\mathbf{x})$ : Modeled in three ways

1. as a single value  $\xi$  for the whole region
2. as separate values  $\xi_{mtn}, \xi_{plains}$
3. as a Gaussian process as above

# Exceedances Model - Priors

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*Priors for  $\alpha_{\phi, \cdot}$ :* Non-informative

$$\alpha_{\phi, \cdot} \sim \text{Unif}(-\infty, \infty)$$

# Exceedances Model - Priors

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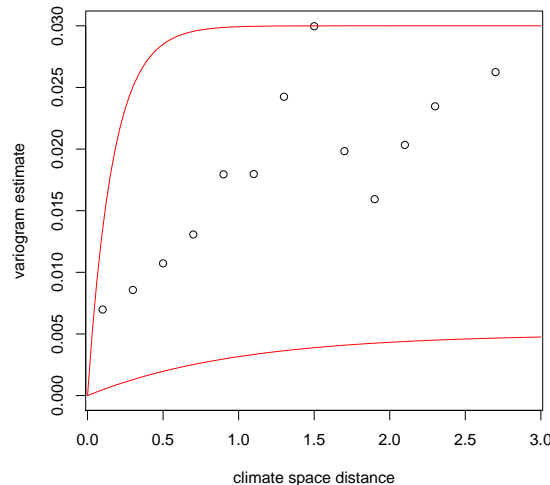
Priors for  $\alpha_{\phi, \cdot}$ : Non-informative

$$\alpha_{\phi, \cdot} \sim \text{Unif}(-\infty, \infty)$$

Priors for  $\beta_{\phi, \cdot}$ : Based on empirical information – difficult to elicit prior information

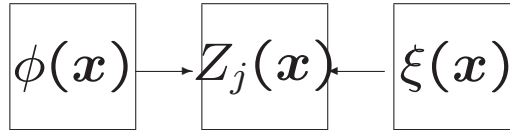
$$\beta_{\phi, 0} \sim \text{Unif}(0.005, 0.03)$$

$$\beta_{\phi, 1} \sim \text{Unif}(1, 6)$$



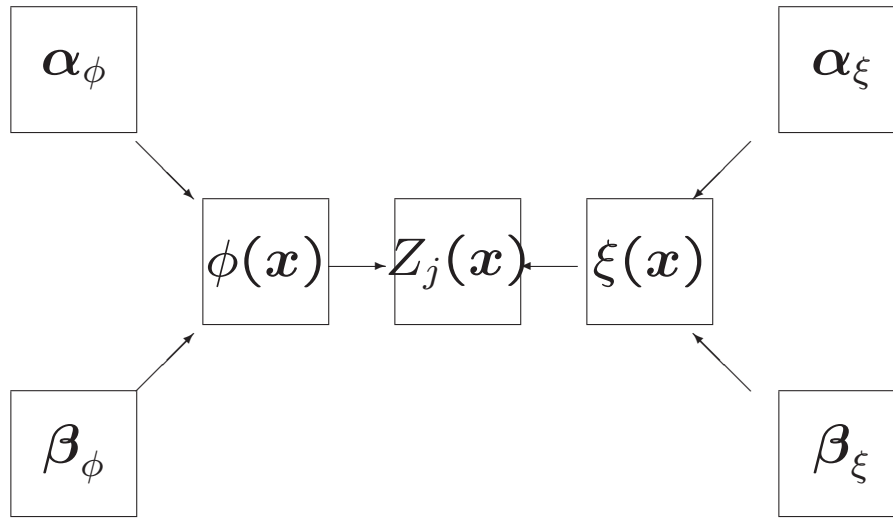
# Model Schematic

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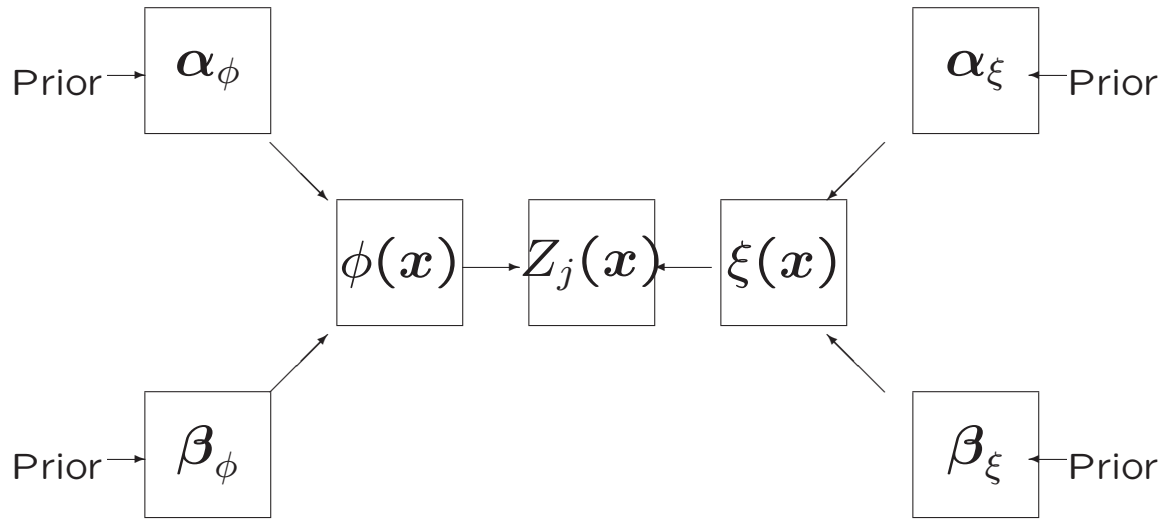
# Model Schematic

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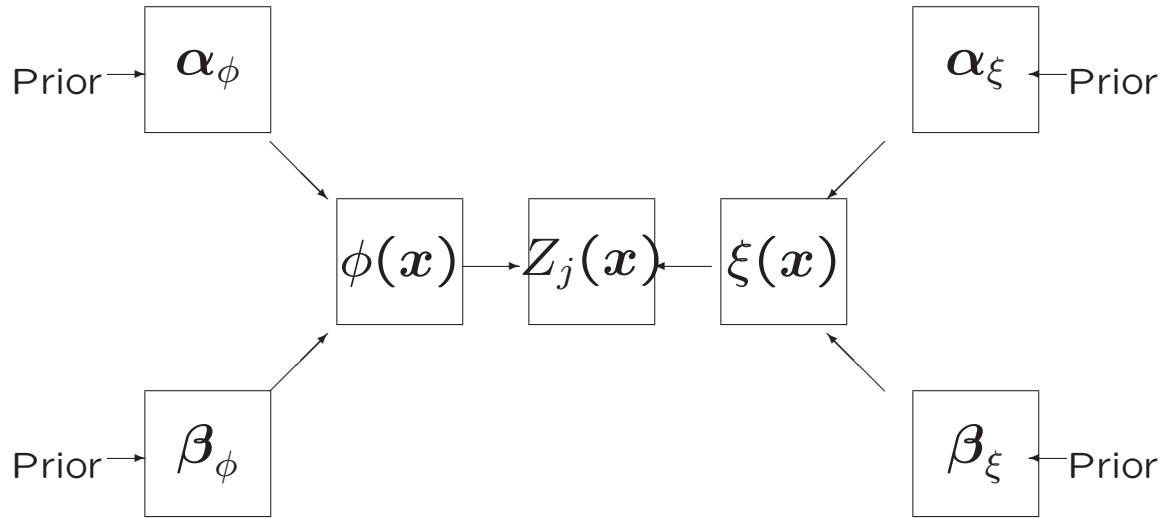
# Model Schematic

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# Model Schematic

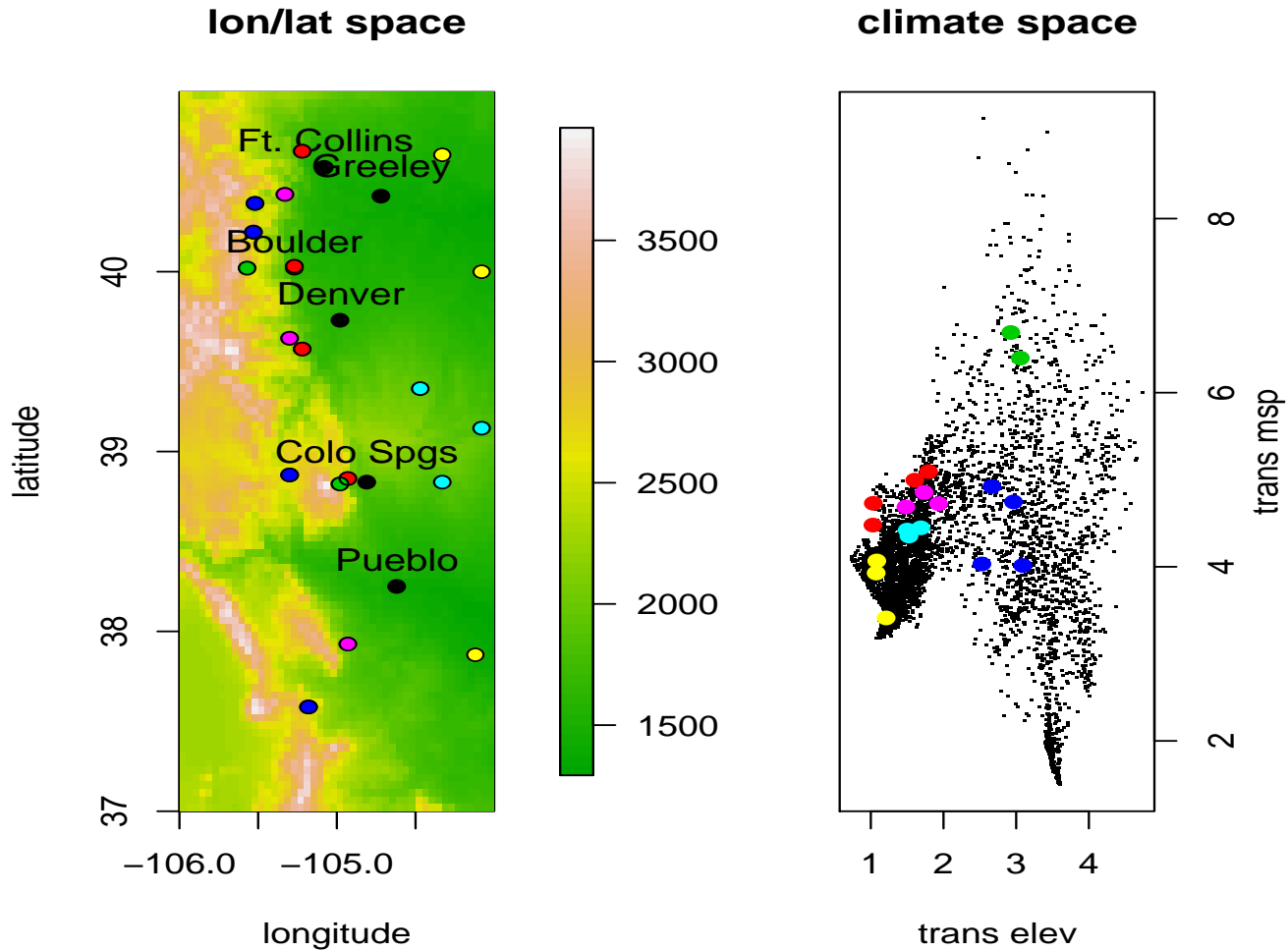
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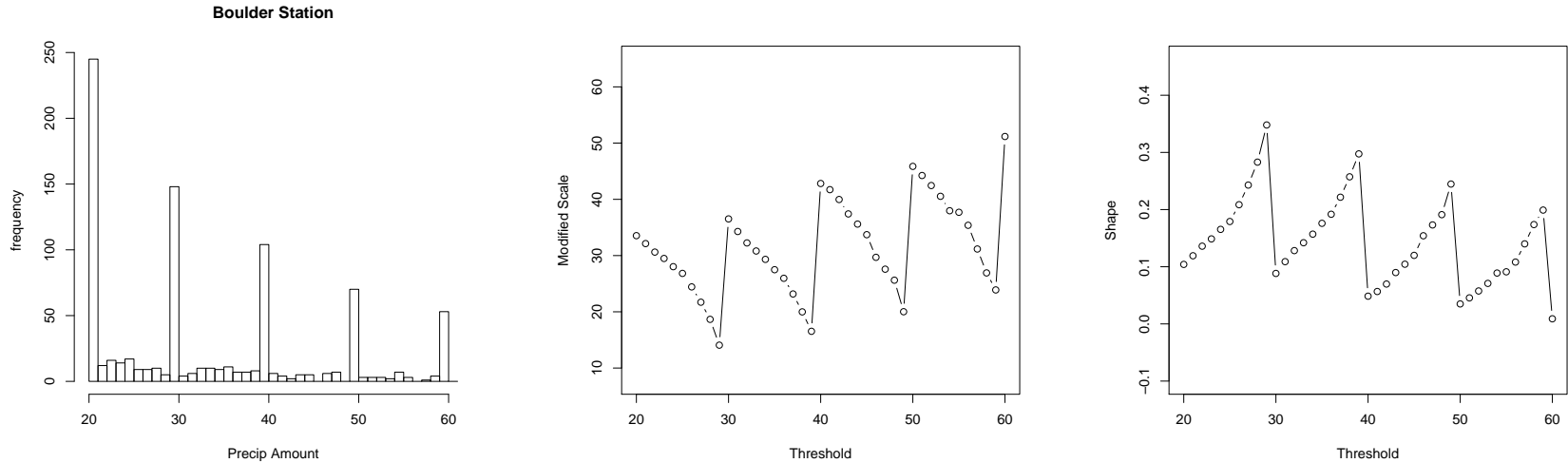
Model assumes that the observations are temporally and spatially independent (conditional on the stations' parameters).

# Climate Space

Problem: Difficult to obtain convergence for  $\beta_{\phi,1}$ .



# Threshold Selection: Quantile or Level?



**Simulation experiment:** Parameter bias least if threshold is chosen in the middle of the precision interval.

Threshold chosen at .45 inches for all stations

⇒  $\zeta_u(\mathbf{x})$  modeled spatially

⇒ Exceedance Rate Model

# Exceedance Rate Model

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*Data Layer:* Assume each station's number of exceedances  $N_i$  is a binomial random variable with  $m_i$  trials each with a probability of  $\zeta(\mathbf{x}_i)$

*Process Layer:* Assume  $\text{logit}(\zeta(\mathbf{x}))$  is a Gaussian process, with mean and covariance

$$\mu_\zeta(\mathbf{x}) = f_\zeta(\boldsymbol{\alpha}_\zeta, \text{covariates}(\mathbf{x}))$$

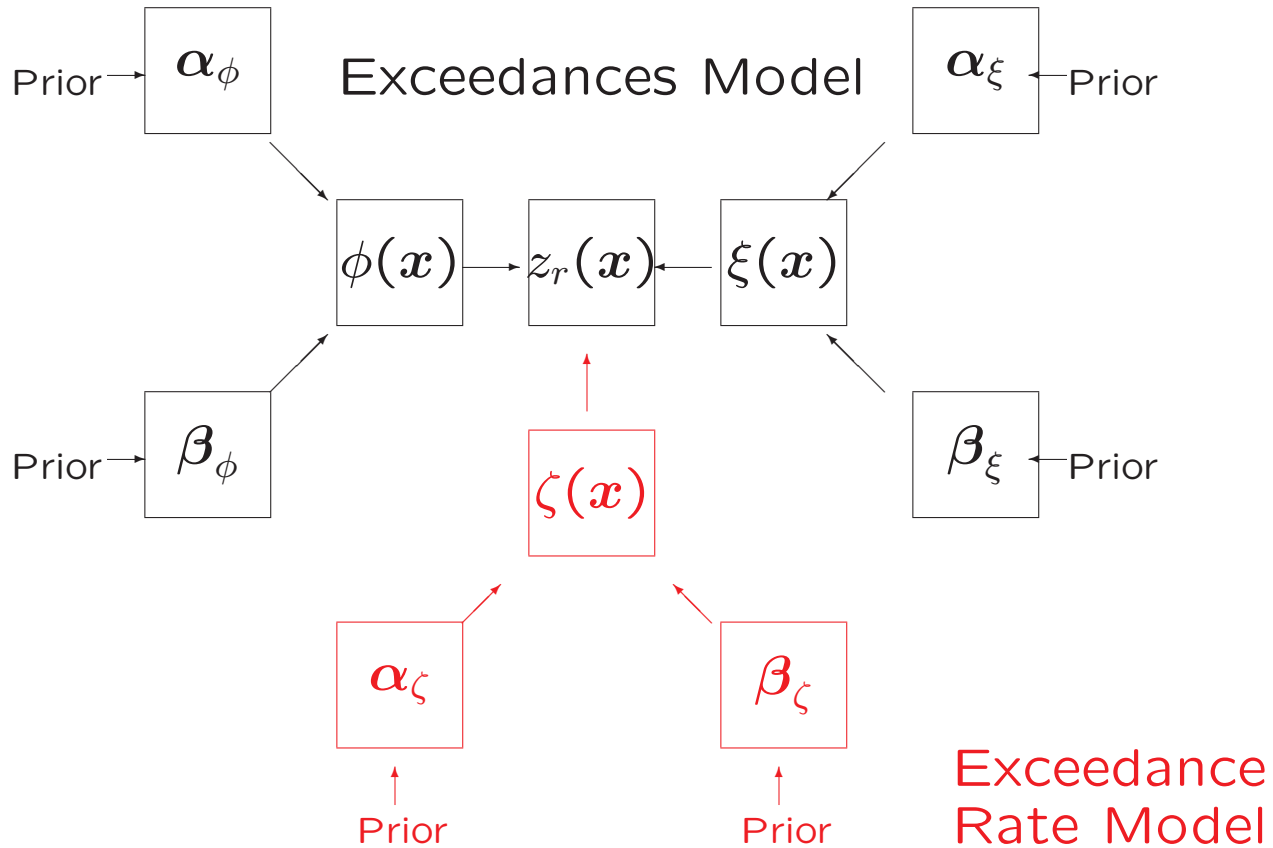
$$\text{Cov}(\zeta(\mathbf{x}), \zeta(\mathbf{x}')) = k_\zeta(\mathbf{x}, \mathbf{x}') = \beta_{\zeta,0} * \exp(-\beta_{\zeta,1} * \|\mathbf{x} - \mathbf{x}'\|_2)$$

*Priors:*

- $\alpha_{\zeta,\cdot} \sim \text{Unif}(-\infty, \infty)$ ,
- $\beta_{\zeta,0} \sim \text{Unif}(0.005, .2)$
- $\beta_{\zeta,1} \sim \text{Unif}(1, 6)$

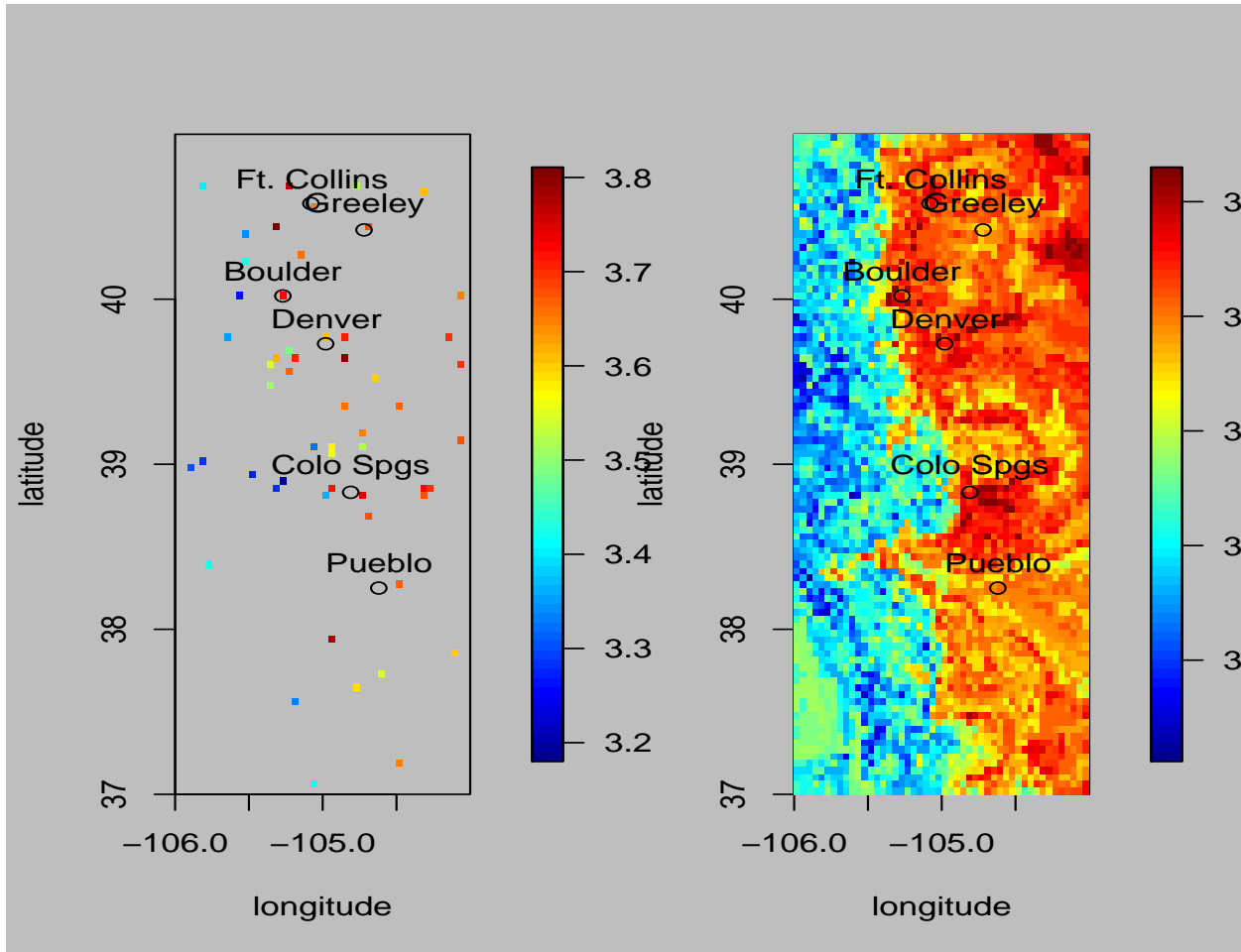
# Model Schematic for Return Levels

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# Interpolation Method

Draw values from  $[\phi(\mathbf{x}) | \phi(x_1), \dots, \phi(x_{56}), \alpha_\phi, \beta_\phi]$ .



# Exceedance Models Tested

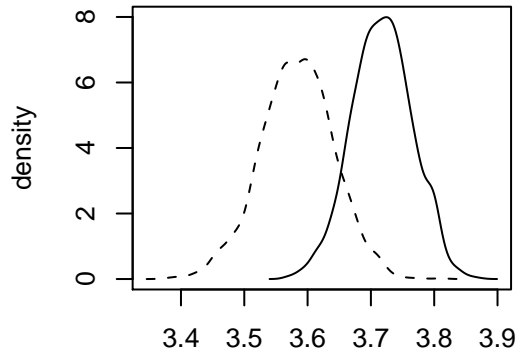
Baseline Model		$\bar{D}$	$p_D$	$DIC$
Model 0:	$\phi = \phi$ $\xi = \xi$	112264.2	2.0	112266.2
Models in Latitude/Longitude Space		$\bar{D}$	$p_D$	$DIC$
Model 1:	$\phi = \phi + \epsilon_\phi$ $\xi = \xi$	98533.2	33.8	98567.0
Model 2:	$\phi = \alpha_0 + \alpha_1(\text{msp}) + \epsilon_\phi$ $\xi = \xi$	98532.3	33.8	98566.1
Model 3:	$\phi = \alpha_0 + \alpha_1(\text{elev}) + \epsilon_\phi$ $\xi = \xi$	98528.8	30.4	98559.2
Model 4:	$\phi = \alpha_0 + \alpha_1(\text{elev}) + \alpha_2(\text{msp}) + \epsilon_\phi$ $\xi = \xi$	98529.7	29.6	98559.6
Models in Climate Space		$\bar{D}$	$p_D$	$DIC$
Model 5:	$\phi = \phi + \epsilon_\phi$ $\xi = \xi$	98524.3	27.3	98551.6
Model 6:	$\phi = \alpha_0 + \alpha_1(\text{elev}) + \epsilon_\phi$ $\xi = \xi$	98526.0	25.8	98551.8
<b>Model 7:</b>	$\phi = \alpha_0 + \alpha_1(\text{elev}) + \epsilon_\phi$ $\xi = \xi_{\text{mtn}}, \xi_{\text{plains}}$	<b>98524.0</b>	<b>26.0</b>	<b>98550.0</b>
Model 8:	$\phi = \alpha_0 + \alpha_1(\text{elev}) + \epsilon_\phi$ $\xi = \xi + \epsilon_\xi$	98518.5	79.9	98598.4

$\epsilon. \sim MVN(0, \Sigma)$  where  $[\sigma]_{i,j} = \beta_{.,0} \exp(-\beta_{.,1} \|x_i - x_j\|)$

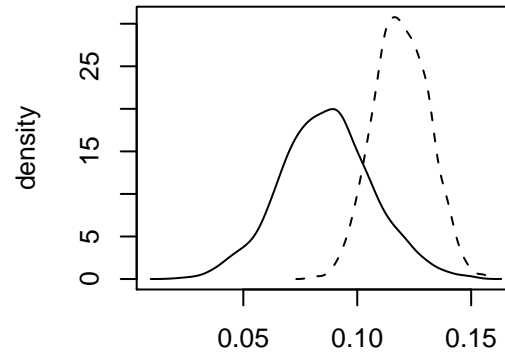
# Posterior Distributions

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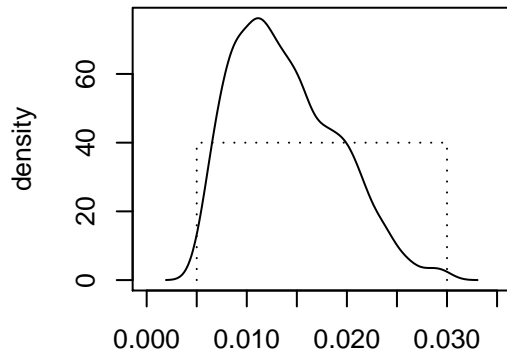
**phi**



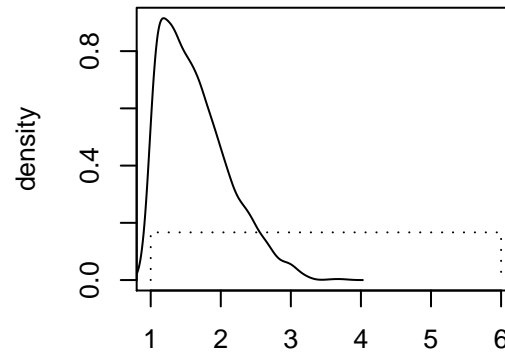
**xi**



**beta\_0 (Sill)**

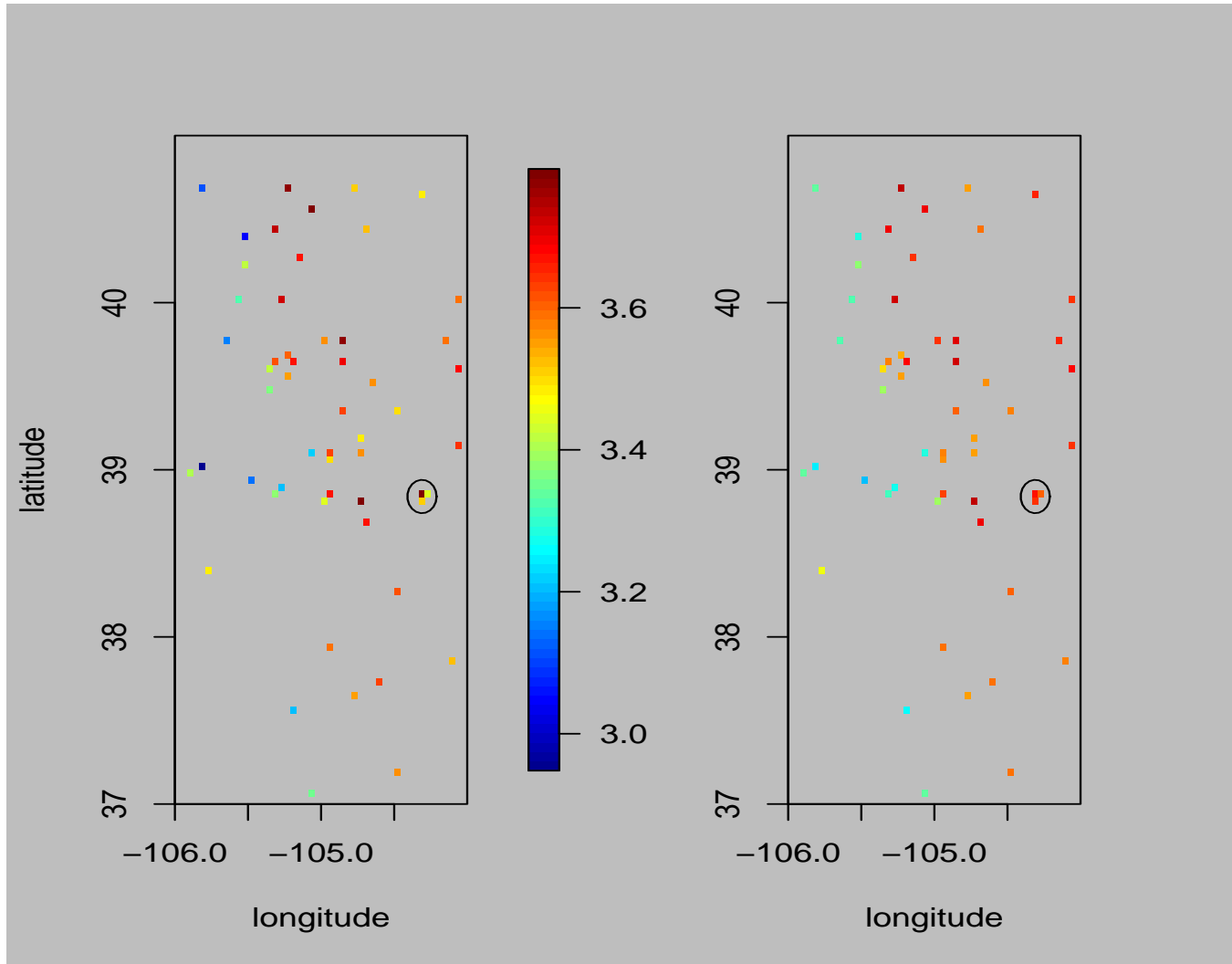


**beta\_1 (Range)**



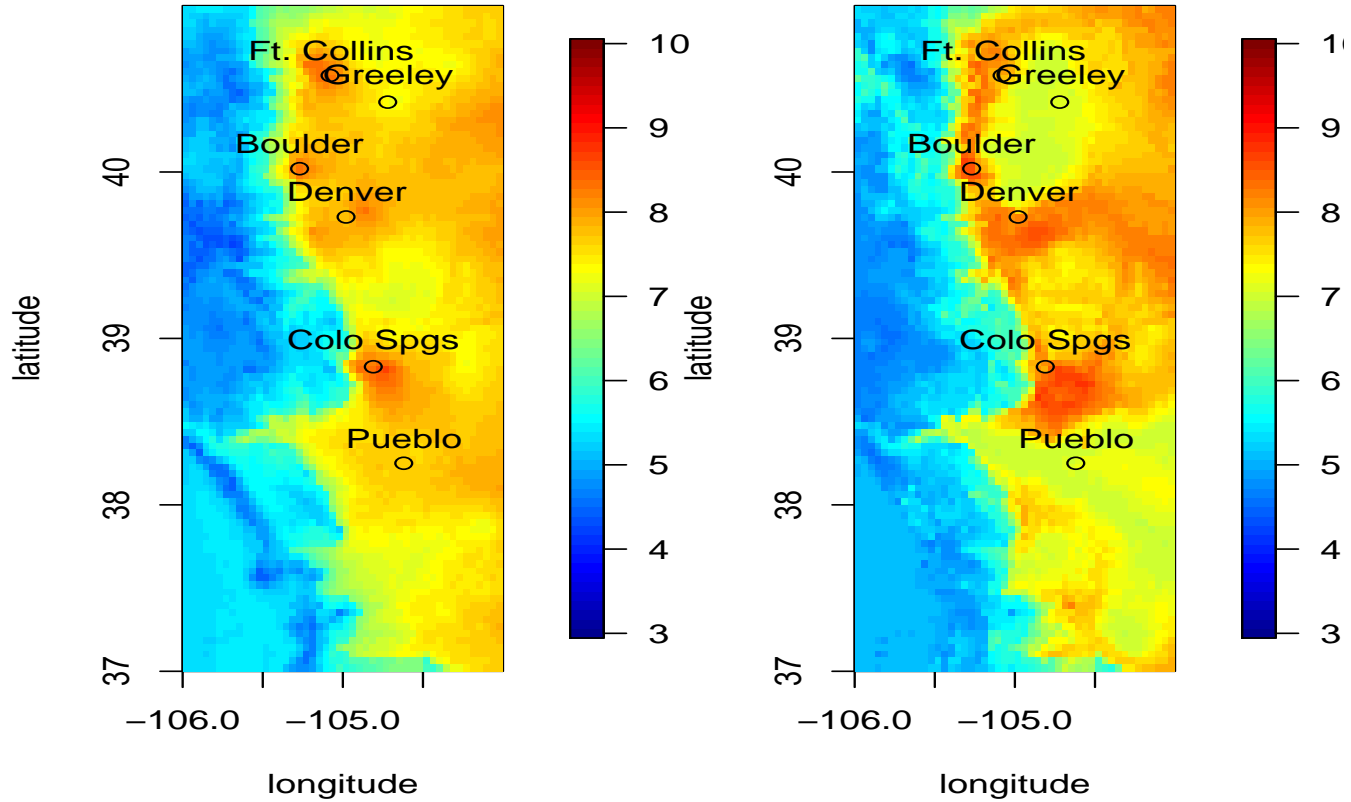
# Spatial Coherence of $\phi$

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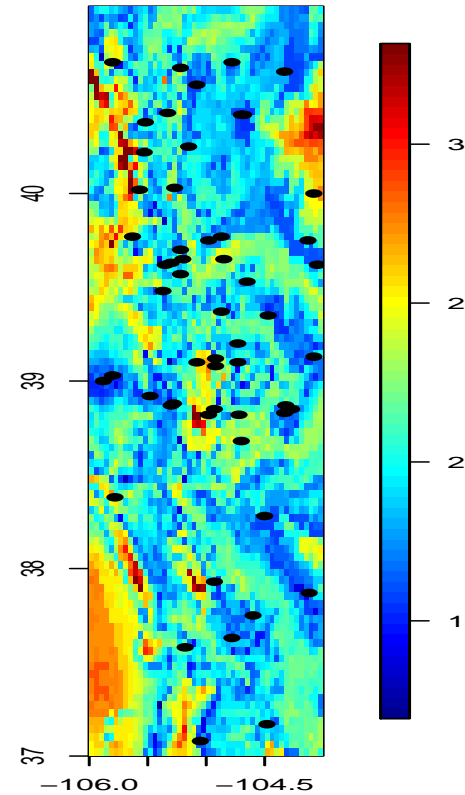
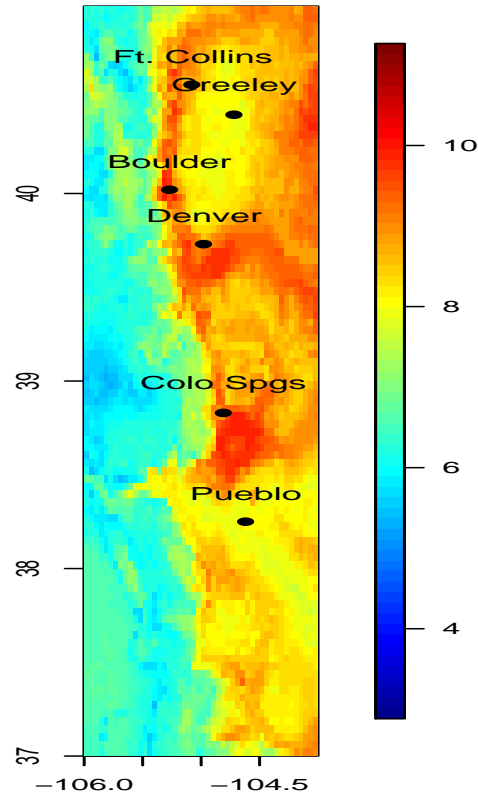
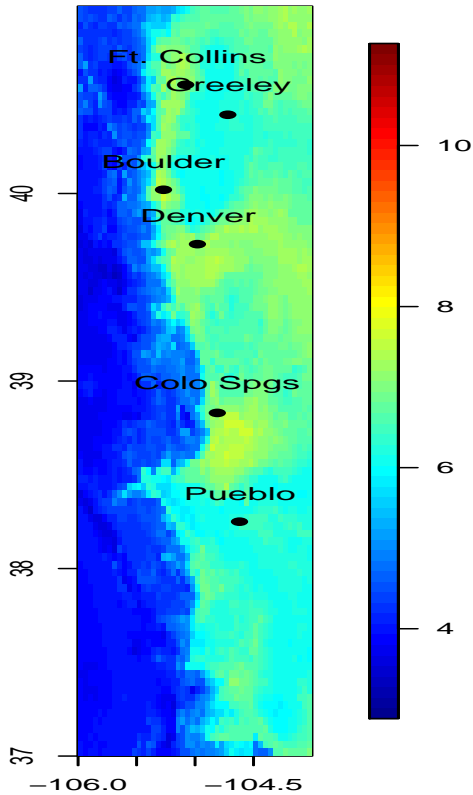
# Traditional vs Climate Space

## 25-year Return Level Point Estimate



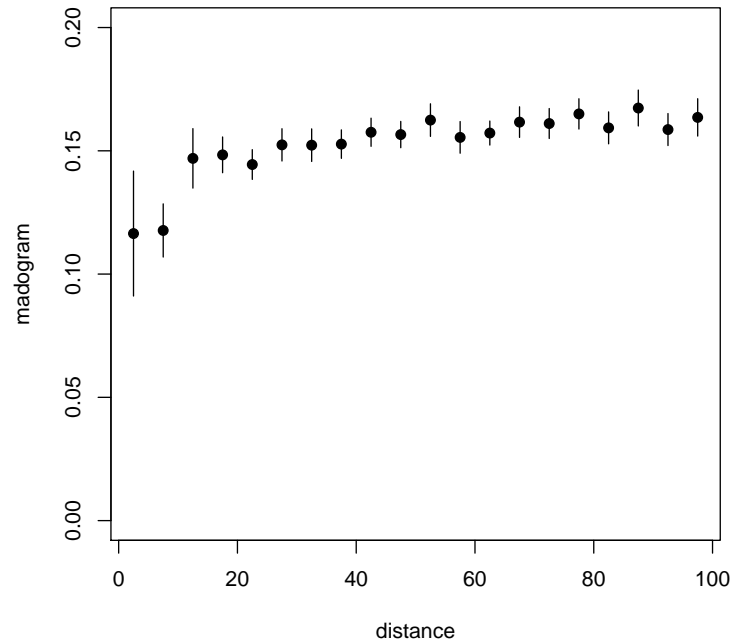
# Results for Model 3: Return Level Uncertainty

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# Madogram on Colorado Data

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- Renormalized *annual max* data.
- Shows very short range dependence in annual max observations.
- Like to apply the madogram to GPD data.

# Conclusions and Future Work

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- Used a Bayesian hierarchical model to produce maps which characterize *climatological* extreme precipitation.
  - Method accounts for uncertainty due to both parameter estimation and interpolation.
  - Dealt with issue of low precision by modeling exceedance rate spatially.
  - Performed spatial analysis in a non-traditional climate space.
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- Extend idea to other data (ozone levels).
  - Model for all duration periods.